

A fiery wake-up call for climate science

To improve climate resilience for extreme fire events, researchers need to translate modelling uncertainties into useful guidance and be wary of overconfidence. If Earth system models do not capture the severity of recent Australian wildfires, development is urgently needed to assess whether we are underestimating fire risk.

Benjamin M. Sanderson and Rosie A. Fisher

The images from the Australian fires served as a bitter climax to a year that was already dominated by climate change and climate-related extreme events. A natural disaster of breathtaking scale unfolds as large fractions of Australia's east coast have burned to an extent not seen in living memory, releasing an estimated 350 million tonnes of CO₂ into the atmosphere in November and December¹, and causing

the loss of thousands of homes and the death of hundreds of millions of animals².

The natural and human disaster has clear political resonance in a country with contentious climate politics, and the question of whether anthropogenic climate change has caused or exacerbated the fires has been a central topic of public debate. Scientists drawn into this debate are generally expected to provide 'hot takes',

which are often distilled into strong attribution statements in the media^{3,4} or explicitly make the case for a simple relationship between warming and fire behaviour⁵⁻⁷. But the relationship between fire and climate is notably complex⁸, and blanket simplifying statements risk undermining expert authority. Worse, by failing to recognize and address knowledge gaps, we may leave society unprepared for

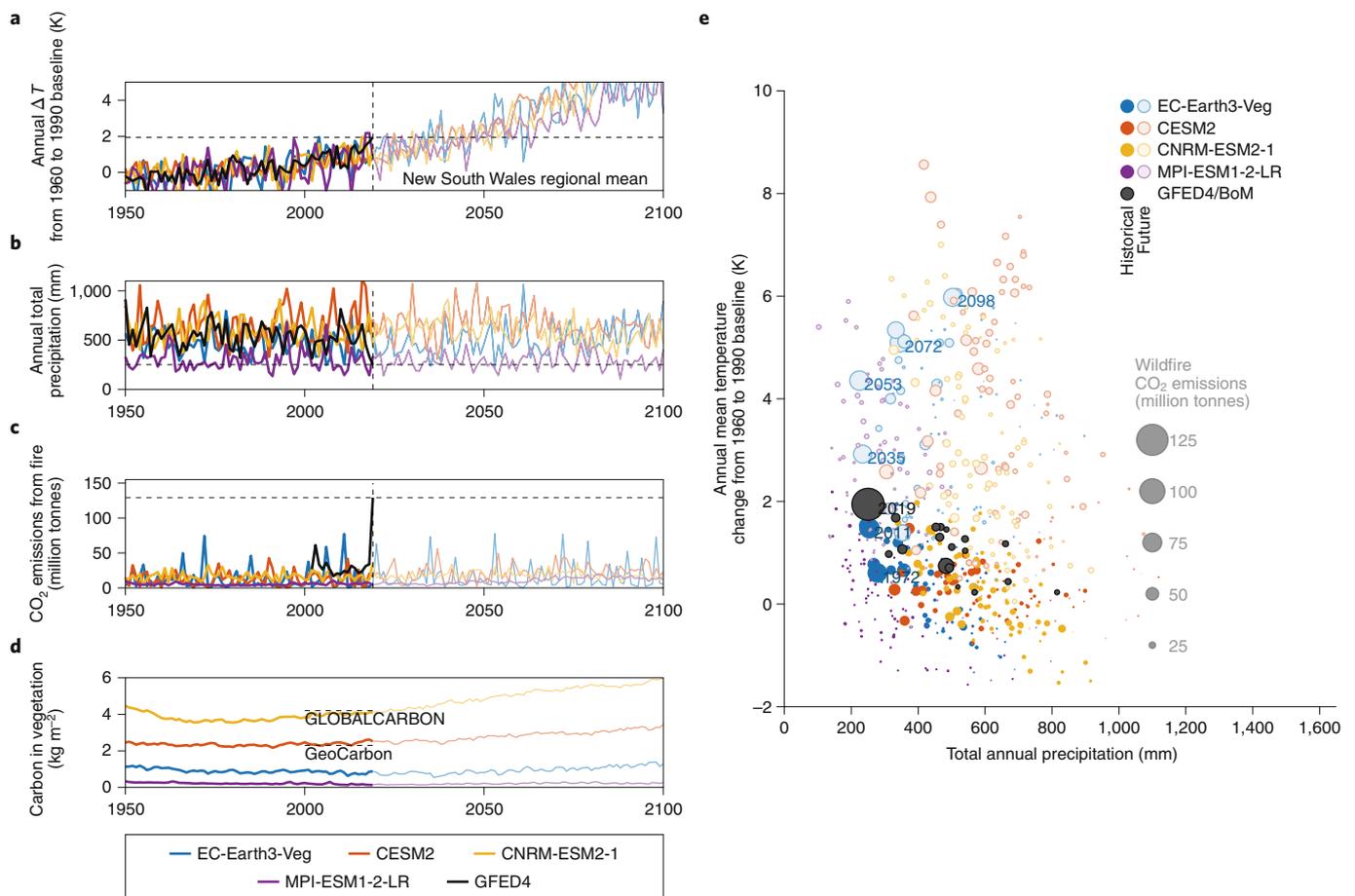


Fig. 1 | Observed and projection of historical and future climate in New South Wales, Australia. Projections are from four CMIP6 models that include fire model output in scenario projections. **a–d**, Evolution of the annual mean surface-temperature anomaly (from a 1960 to 1990 baseline); absolute precipitation; CO₂ emissions from fire; and vegetation biomass. **e**, Fire CO₂ emissions (illustrated by bubble size) as a function of both surface temperature anomaly and total annual mean precipitation. Future projections are shown in faded colours. Mean observed biomass estimates are from GLOBALCARBON³⁶ and GeoCarbon³⁷. Temperature and precipitation data are from the Australian Bureau of Meteorology³⁸. Fire emissions data are from the Global Fire Emissions Database (GFED)¹.

potentially more extreme events in the future.

For an increasing number of catastrophes, strong attribution statements are justified. Mean warming levels are now sufficiently large that many high-temperature extreme events would be impossible without anthropogenic influence⁹, and they can be reliably projected to become more intense in the future. In the case of recent events in Australia, there is no doubt that the record temperatures of the past year would not be possible without anthropogenic influence, and that under a scenario where emissions continue to grow, such a year would be average by 2040 and exceptionally cool by 2060 (Fig. 1a).

If all else is kept constant, higher temperatures will result in more fire-prone conditions¹⁰. This is represented in operational fire risk metrics used in forest management, which are calculated as functions of temperature, wind, moisture and fuel availability¹¹. Such metrics, however, are calculated on historical datasets, and a premise that these relationships will hold in future climate is an extrapolation¹².

The complex dynamics of fuel accumulation, vegetation dynamics and their interactions with climate under transient CO₂ concentrations, as well as impacts of land management and human ignitions, are likely to result in fire behaviour patterns not represented in historical records. Thus, fire prediction over decadal to century timescales requires more mechanistic approaches, capable of capturing the numerous interacting system components that affect the evolution of fire risk^{13,14}.

Process-based global fire models based on these principles have progressed rapidly over the past decade^{15–17}. Their use in fully coupled climate projections, however, is still not standard practice. Many Earth system models are thus omitting a potentially important component of the global carbon-climate feedback, while failing to deliver projections of one of the main facets of climate impacts on human society^{18,19}. For those few CMIP6 models that do include prognostic fire (four unique models with future coupled projections: EC-Earth3-Veg, CESM2, CNRM-ESM2-1 and MPI-ESM1-2-LR), there can be large regional biases that make projections or formal attribution statements difficult.

For the case of New South Wales (the Australian state in which the fire extent is most unprecedented), the scale of the fires is unmatched in the CMIP6 simulations in either the present or the future (Fig. 1c; similar biases are evident in other Australian territories, as illustrated in the Supplementary Information). This is partly because 2019

was marked also by exceptionally low rainfall (Fig. 1b). For EC-Earth3-Veg (the one CMIP6 model that does simulate occasional mega-fires in southeast Australia), rainfall is a much stronger predictor of fire extent than temperature. Indeed, in that model, it is only at much greater regional warming levels — 4 °C above pre-industrial — that similar fire extents are seen in years without comparably low rainfall (Fig. 1e). Precipitation projections for southeast Australia remain highly uncertain and model-dependent¹⁰, and assessing the changing probabilities of low-rainfall years like 2019 under climate change requires large ensembles that simulate many realizations of natural variability²⁰.

Fundamentally, modelling the cascade of climate, vegetation and anthropogenic feedbacks that lead to extreme fire events in semi-arid regions is challenging. In coupled Earth system models, biases in climate (or productivity) can mean that some models do not simulate enough vegetation to allow large burns to occur (for example MPI — purple in Fig. 1). Conversely, in systems that are wet or dominated by woody plants, vegetation- or fuel-mediated feedbacks may act to maintain systems in a fire-free state (both in models²¹ and real life^{22,23}). Simulation of transitional semi-arid vegetation is in itself difficult²⁴ on account of poorly understood phenology^{25,26}, root water access^{27,28}, grass/understorey dynamics²⁹, fuel dynamics⁸ and heterogeneity. Fire ignition also requires parameterization and calibration. Fire models typically initialize burning events through both lightning-induced and anthropogenic ignitions, the latter based on a probabilistic function of human population density (as is the effectiveness of fire suppression activities)³⁰.

Recently published results of the first 'FireMIP' intercomparison project^{17,31} have illustrated the capabilities of global fire models (and their host land-surface schemes) to capture these interactions when driven 'offline' with climate reanalysis data. They shed light on the causes of variability in model responses to climate and vegetation state⁸, as well as illustrating the importance of appropriate representation of land use and human ignitions³⁰. If we aspire to make useful projections of the future risk of catastrophic fires, entraining this new understanding into coupled Earth system simulations must be a high priority of upcoming model development efforts. At present, the inclusion of fire is arguably considered as an afterthought (or not at all) by many Earth system modelling efforts, representing a mismatch between resources dedicated to understanding this problem and the seriousness of its potential consequences.

It is critical that Earth system modelling is capable of informing the changing risk of potentially devastating events. The fact that Australia has experienced damages that go beyond what is currently simulated highlights that current syntheses may be missing major risks. Policymakers should take this as a warning that ongoing emissions will take us into an increasingly unpredictable climate space where impacts may be more extreme than projections. Scientists, on the other hand, need to tread a delicate line of underlining what is certain and providing appropriate guidance on what is not, while redoubling efforts to better represent climate impacts that most directly affect society. □

Benjamin M. Sanderson^{1,2}  and Rosie A. Fisher¹

¹CERFACS, Toulouse, France. ²National Center for Atmospheric Research, Boulder, CO, USA.

✉e-mail: sanderson@cerfacs.fr

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References

- Giglio, L., Randerson, J. T. & van der Werf, G. R. *J. Geophys. Res. Biogeosci.* **118**, 317–328 (2013).
- University of Sydney News (8 January 2020); <https://go.nature.com/3aBhlyu>
- Eggleton, M. *National Geographic* (15 November 2019); <https://go.nature.com/38ME5I>
- Rice, D. *USA Today* (8 January 2020); <https://go.nature.com/36xAvfA>
- Mann, M. *The Guardian* (1 January 2020); <https://go.nature.com/30WmreF>
- Law, T. *Time* (7 January 2020); <https://go.nature.com/2RwgBxx>
- Abram, N. *Scientific American Blog Network* <https://go.nature.com/3aNF16o> (2019).
- Forkel, M. et al. *Biogeosciences* **16**, 57–76 (2019).
- Sippel, S., Meinschausen, N., Fischer, E. M., Székely, E. & Knutti, R. *Nat. Clim. Change* **10**, 35–41 (2020).
- Kirchmeier-Young, M. C., Zwiers, F. W., Gillett, N. P. & Cannon, A. J. *Climatic Change* **144**, 365–379 (2017).
- Van Wagner, C. E. & Canadian Forestry Service. *Development and Structure of the Canadian Forest Fire Weather Index System*. Forestry Technical Report 35 (Canadian Forestry Service, 1987).
- Clarke, H. & Evans, J. P. *Theor. Appl. Climatol.* **136**, 513–527 (2018).
- Hoffmann, W. A. et al. *Austr. Ecol.* **37**, 634–643 (2012).
- Shuman, J. K. et al. *Environ. Res. Lett.* **12**, 035003 (2017).
- Thonicke, K. et al. *Biogeosciences* **7**, 1991–2011 (2010).
- Hantson, S. *Biogeosciences* **13**, 3359–3375 (2016).
- Hantson, S. et al. *Geosci. Model Dev.* <https://doi.org/10.5194/gmd-2019-261> (2020).
- Liu, Y., Goodrick, S. & Heilman, W. *For. Ecol. Manag.* **317**, 80–96 (2014).
- Arora, V. K. et al. *Biogeosci. Discuss.* <https://doi.org/10.5194/bg-2019-473> (2019).
- Norris, J., Chen, G. & Neelin, J. D. *J. Clim.* **32**, 5397–5416 (2019).
- Scheiter, S., Moncrieff, G. R., Pfeiffer, M. & Higgins, S. I. *Biogeosci. Discuss.* <https://doi.org/10.5194/bg-2019-415> (2019).
- Staver, A. C., Archibald, S. & Levin, S. A. *Science* **334**, 230–232 (2011).
- Forkel, M. et al. *Environ. Res. Commun.* **1**, 051005 (2019).
- Whitley, R. et al. *Biogeosciences* **14**, 4711–4732 (2017).
- Dahlin, K. M., Fisher, R. A. & Lawrence, P. J. *Biogeosciences* **12**, 5061–5074 (2015).
- Dahlin, K. M., Ponte, D. D., Setlock, E. & Nagelkirch, R. *Ecography* <https://doi.org/10.1111/ecog.02443> (2017).
- Williams, M., Law, B. E., Anthoni, P. M. & Unsworth, M. H. *Tree Physiol.* **21**, 287–298 (2001).
- Kauwe, M. G. D. et al. *Biogeosciences* **12**, 7503–7518 (2015).
- Whitley, R. et al. *Biogeosciences* **13**, 3245–3265 (2016).
- Teckentrup, L. et al. *Biogeosciences* **16**, 3883–3910 (2019).
- Rabin, S. S. et al. *Geosci. Model Dev.* **10**, 1175–1197 (2017).

32. Neubauer, D. et al. HAMMOZ-Consortium MPI-ESM1.2-HAM model output prepared for CMIP6. <https://doi.org/10.22033/ESGF/CMIP6.5016> (2019).
33. Seferian, R. CNRM-CERFACS CNRM-ESM2-1 model output prepared for CMIP6 CMIP. <https://doi.org/10.22033/ESGF/CMIP6.1391> (2018).
34. EC-Earth Consortium (EC-Earth). EC-Earth-Consortium EC-Earth3-Veg model output prepared for CMIP6 ScenarioMIP. <https://doi.org/10.22033/ESGF/CMIP6.727> (2019).
35. Danabasoglu, G. NCAR CESM2 model output prepared for CMIP6 ScenarioMIP. <https://doi.org/10.22033/ESGF/CMIP6.7768> (2019).
36. Liu, Y. Y. et al. *Nat. Clim. Change* **5**, 470–474 (2015).
37. Ge, Y., Avitabile, V., Heuvelink, G. B. M., Wang, J. & Herold, M. *Int. J. Appl. Earth Obs.* **31**, 13–24 (2014).
38. Australian Government Bureau of Meteorology. Climate change — trends and extremes. <http://www.bom.gov.au/climate/change> (accessed 8 January 2020).

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Additional information

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